OCCURRENCE OF SOME MAASTRICHTIAN DINOFLAGELLATE CYSTS FROM THE UPPER CRETACEOUS SEDIMENTS OF THE BIDA BASIN, NIGERIA: IMPLICATIONS FOR AGE AND PALEOENVIRONMENTS.

J. OJO, OLUSOLA

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ABSTRACT

Lithostratigraphic and palynological studies of the shale facies of the upper Cretaceous Patti Formation, southeastern Bida Basin, Nigeria have allowed the determination of the paleoenvironment and age of the sediments. Marine dinocyst assemblage and terrestrial pollen and spores (relatively more abundant) are well preserved in the sections. The marine dinocysts include diverse species of the genera Dinogymnium, Senegalinium, Andalusialla and Spiniferites. This assemblage reflects influence of shallow marine process during sedimentation of the shales and they provide, for the first time, incontrovertible evidence of the presence of marine condition in southeastern part of the Bida Basin.

A Maastrichtian age is suggested for the Shale facies of the Patti Formation in the study area based on the occurrence of well known Maastrichtian marine dinoflagellates Dinogymnium acuminatum, Dinogymnium digitus, Senegalinium bicavatum, Senegalinium psilatum and Paleocystodinium australinium. Associated with these forms also are some Maastrichtian spores and pollen; Buttinia andreevi, Cristaeturites cristas, and Retidiporites magdalenensis.

This study therefore, is considered significant as it also confirms marginal marine connection between the Anambra and Bida basins and consequently is a proof that the later formed a part of the links between the Tethys sea and the south Atlantic sea in the Maastrichtian time.

KEY WORDS: Bida Basin, Patti Formation, Maastrichtian

INTRODUCTION

The Bida Basin also known as middle Niger Basin or Nupe Basin lies within west central Nigeria and trends in the northwest southeast direction (Fig. 1). This hinterland Basin consists of Upper Cretaceous clastic sequence of sandstones, shales, siltstones and ironstones which are considered non petrolierous. Previous account of the geology indicates a rift related origin (King 1950; Kogbe et al., 1981; Ojo and Ajakaiye, 1989), associated with the Benue Trough system and drifting apart of African and Brazilian plates.

The Patti Formation which is the only stratigraphic unit containing carbonaceous shales in the Bida Basin is sandwiched between the older Campanian - \ Maastrichtian Lokoja Formation (conglomerates, sandstones and claystones) and younger Agbaja Formation of mainly ironstone (Fig 2). Several studies particularly on sedimentology and lithostratigraphy of the sediments and their lateral equivalents in the northwest Bida basin have been undertaken and reported (Jones, 1958; Adeleye, 1973, 1974; Braide, 1992; Ladipo et al., 1994; Abimbola et al., 1999; Vrbkha et al., 1999). Recently, Ojo and Akande (2003) and Akande et al. (2005) provided detail information and interpretation of the sedimentological and organic geochemical characteristics of the sediments. Notably, published works dealing with biostratigraphy in the present study area are scarce. The few available ones are those of Jan Du Chene et al. (1978), Mebradu et al. (1986) and Agyingi (1993). The impression from these studies, which reported only non marine pollen/spores from the shales of the Patti Formation, is that the shales and siltstones are non - marine. The recent work of Ojo and Akande (2008) however indicated the presence of marine dinoflagellates in some intervals of the Patti Formation sections. The aims of this paper are therefore to i) describe the dinoflagellate assemblage recovered from the study area ii) highlight the biostratigraphic significance of this assemblage and provide evidence of marine environment in the southeastern Bida Basin. This is important because it will help to correct the wrong but long held impression of zero marine condition in the study area. The new data is also expected to provide more understanding of the paleogeography of the Bida basin with respect to possible connection between the Tethys and the south Atlantic in the Late Cretaceous and also may also help resolve more resolutely the precise age of the Patti sediments.
Due to lack of exploration wells or well preserved borehole samples in the study area, the samples, mainly shale facies were obtained from road and quarry exposures of the Patti Formation which are restricted to the centre of the Basin around Ahoko, along Lokoja – Abuja highway (Fig. 1). Two lithologic sections (quarry face and road cut) representing the shale – claystone member of the Patti Formation (Ojo and Akande, 2006) were investigated on the basis of bed to bed measurements and systemic sampling of relevant intervals. The road cut section (about 26m thick) consists of well bedded, alternating shales, siltstones, claystones and ironstones (Fig. 3). At the lower part of the section are beds of fissile, carbonaceous shale with average thickness of 0.5m which exude sulfurous odour. The quarry section is about 20m thick and the lithologies are similar to those in the road section (Fig. 4). Effort was made to reach the deeper part of the shales to avoid weathered samples. In both sections, 40 samples of the fresh carbonaceous shales and siltstones were selected and used in this study for palynological preparation. The palynological technique involved standard chemical extraction or maceration with application of 10% HCl to remove the carbonates and 40% HF for the silicate digestion. The heavy minerals were separated from the residue using heavy liquid (Na – polywolfromat, 2.0 S.G). The separates were screened through 10μm nylon sieve. Due to high carbon content of most of the samples, ultrasonic machine was used to clean the residues. Clean palynomorphs were subsequently mounted on glass slides using glycerine jelly. Microscopic study involving identification, counting and photographic documentation were done using Leitz Diaplan microscope.

Dinoflagellate Assemblage

The dinoflagellate cyst assemblage, which are restricted only to the lower part of the studied sections are relatively abundant but of low diversity. The microplanktons are mainly represented by the peridinoids which account for 50% (Fig 5). Others are spinerates (33%), microforaminiferal lining (13.5%) and algae (3.5%).

Description of some figured dinoflagellates and spores/pollen

Senegalinium bicavatum (Plate 1 Fig. 2)
Description: Shell is ovate shaped and non tabulate, intercalary archeopyle, tranverse furrow not distinct and longitudinal furrow present. One broad but short apical horn, 2 antapical horns of near equal lengths. Thin periphram but endophragm is relatively thick.
Remarks: This species has been recorded in Senegal Basin (Jain and Miliepied, 1975), Maastrichtian Senegalianium of Egypt (Schrank, 1984).

Senegalinium psilatum (Plate 1 Fig. 4)
Description: Bicavate, shell is ovate, intercalary archeopyle below the apical horn, well developed tranverse furrow. Very short apical and antapical horns. Periphram is delicate and psilate.
Remarks: Well known and reported in Campanian – Maastrichtian of Senegal (Jain and Miliepied, 1975), Senonian to Maastrichtian, Gabon (Malloy, 1972) and Maastrichtian Gombe Formation, northeastern Nigeria (Ojo and Akande, 2004).

Dinogymnium euclaensis (Plate 1 Fig. 6)
Description: Shell is biconical, cingulum prominent and divides shell into unequal halves, epicyst is longer than the hypocyst, apical end is blunt and more rounded than the antapical end. Archeopyle is apical. Sulcus is broad and distinct.
Remarks: This species has been recorded in Maastrichtian S.E. Nigeria (Oloto, 1989), Maastrichtian Senegal (Jain and Miliepied, 1975), Campanian to Maastrichtian sediments of Egypt (Schrank, 1984).

Dinogymnium acumatum (Plate 1 Fig. 7)
Description: Shell is biconical and elongated. Cingulum is prominent, epitract is narrower than the broad hypotrac and tranversed by thin longitudinal folds. Archeopyle is apical. Surface is granular.

Dinogymnium digitus (Plate 1 Fig. 8)
Description: The shell is oblong, epitheca and hypotheeca are almost equal. Archeopyle is apical. Tranverse furrow is distinct but not broad or deep. Longitudinal folds are relatively broad in the epitheca and hypotheeca. Surface is granular.
Remarks: Previously known occurrence include Turonian, Kazakhastan (Vozhzhennikova, 1967).

Paleocystodinium australinium (Plate 1 Fig. 9)
Description: Long and moderately broad, smooth apical and antapical horns, few longitudinal folds. Periphram is close to endophragm.
Remarks: Previous known occurrence include; Maastrichtian of Senegal (Jain and Miliepied, 1973), Maastrichtian – Paleocene of Gabon (Malloy, 1972) and Maastrichtian Nkporo Shale, southern Nigeria (Oloto, 1989).

Retidiporites magdalenensis (Plate 1 Fig. 18)
Description: Shape is cylindrical and having microfoveolate to reticulate sculptural elements.
Remarks: Previously known and reported from Maastrichtian Egypt, Upper Campanian – Maastrichtian Senegal (Schrank, 1987), in the Maastrichtian of southern Nigeria (Salami, 1990).

Cristaturnites cristatus (Plate 1 Fig. 16)
Description: Shape is oval and bilateral symmetry. Inaperturate and sculpture is echinaten
Remarks: It has been recorded in the Maastrichtian W. Africa and Somalia (Schrank, 1994).

Buttinia andreevi (Plate 1 Fig. 17)
Description: Sphaeroidal pollen with smooth exine. It is periporate and tectate.
Remarks:widely reported from Maastrichtian West Africa e.g Gabon. Senegal. Anoala Ivory Coast. Todo. Niger

OCCURRENCE OF SOME MAASTRICHTIAN DINOFLAGELLATE CYSTS

BIOSTRATIGRAPHIC IMPLICATIONS

The present study reveals biostratigraphic significant dinocysts. The main constituents of this assemblage are Dinogymnium euclaensis, Dinogymnium acumatum, Dinogymnium digitus,
Senegalinium bicavatum, Senegalinium psilatum and Paleocystodinium australinum. This assemblage obviously compares closely with well known Upper Cretaceous assemblages in different part of the world and most strikingly Campanian - Maaschrichtian sediments of West Africa. The good representation of the genus Dinogymnium in this section suggests a Maaschrichtian age. Jain et al. (1975) and Evitt (1967) reported D. accuminatum and other Dinogymnium species from Jadukata and Mahadek Formations, India and dated it Maaschrichtian. D. spp such as D. accuminatum and D. Euclaensis have been consistently reported as having marked representation in the Campanian to Maaschrichtian sediments (Schrank, 1984, 1987; Sultan, 1985), Senegal Basin (Jain and Millepied, 1973), Maaschrichtian of Carlife America USA (Evitt, 1967), Upper Cretaceous Brazil (Herngreen, 1975). A number of other Campanian – Maaschrichtian forms also occur in the samples. Notably, Senegalinium bicavatum, Senegalinium psilatum which have been reported from the Maaschrichtian of Senegal, N. W. Africa (Jain and Millepied, 1975) are prominent in the study area. The above comparison suggests a Maaschrichtian age for the sediments. The co - occurrence of D. euclaensis, S. bicavatum and P. australinum provide unimpeachable evidence for age neither younger nor older than Maaschrichtian. Moreover, the Maaschrichtian age appears most reasonable because the well known Tertiary elements (both dinocyst and spores/pollen) are absent. Comparison of this assemblage with those reported from Maaschrichtian Nkporo Shale, Anambra Basin (Oloto, 1989), Gombe Formation, Gongola Basin (Ojo and Akande, 2004) shows striking similarities and positive correlation. It should be noted that the associated pollen and spores (Buttinia andreevi, Retidiporites magdalenensis) from same interval also indicate Maaschrichtian age (Ojo and Akande, 2008).

Paleoenvironmental Synthesis

The environments of deposition of the sediments of Patti Formation have been subject of discussion in various publications (Braide, 1992; Agyingi, 1991). The authors suggested non marine environments for the argillaceous rocks (shale – clayey member) of the Patti Formation. Recent investigation by Ojo and Akande (2006) and Akande et al. (2005) however recognized and reported certain marine sedimentological features such as herringbone and hummocky cross beddings and ophiomorpha burrows in these sediments. The occurrence of well preserved marine dinoflagellate cysts from the shaly intervals of the Patti Formation in the study area, being reported here for the first time, is therefore significant in the elucidation of the processes of sedimentation.

Palynological data have been found useful as paleoenvironmental synthesis tool (Van Bergen et al. 1990; Olot, 1989; Petters and Edet, 1996; Vadja - Santinavez, 1998; Ojo, 1999). In these studies, the interpretation is based on the ratio of land derived miospores to marine dinoflagellate and also the morphology of the dinocysts. Schrank (1984) suggested that a palynomorph assemblage with higher content of land derived miospores indicates terrestrial influence and vice versa. In this study, the occurrence of known marine dinoflagellates such as Dinogymnium, Senegalinium, Spiniferites and microforaminifera lining indicate marine flooding at the lower part of the sections. It is evident, however, from the low diversity of dinocysts with an abundant peridinoids than the chorate cysts that the salinity of the water was low. Harland (1978) posited that the ratio of gonyaulaceans to peridinoids provides an indication of salinity; the higher the ratio, the more open marine condition and vice versa. Here the ratio is extremely low, therefore indicating near shore or brackish water, probably less than 50m (Figs 6&7). The predominance of the peridinoids clearly indicates estuarine environment (Downie et al. 1971), commonly inundated and diluted by run – off water. The unique feature of the Dinogymnium which is well represented in the assemblage from the studied section, is known to be specially adapted to extreme fluctuation in shallow marine settings (May, 1977).

An examination of the chitinous foraminiferal test linings which are common in the samples show that they are single chambered and uniserial. This shows an affinity to bentonthic foraminiferal and therefore an indication of low salinity, brackish water environments. It is notable that the present dinocyst assemblage shows similarities with some of the Campanian – Maaschrichtian dinocysts that are typical of the Tethyan regions. Many of the species are similar to those described from the Upper Cretaceous of southeast and southwest Nigeria (Salami, 1983, 1984). It can be deduced therefore that the Bida Basin provided a link between the Tethyan sea and the Gulf of Guinea in the Maaschrichtian times. The population and diversity of the dinocyst reduced upsection. In deed from the middle to top of section, no marine dinocyst is present. This may be an indication, that the lower part represents regression at the end of Maaschrichtian.
**LEGEND**

- Alluvium Recent and Swamp Forest
- Maastrichtian Agbaja Ironstone Formation
- Maastrichtian Patti Formation
- Campanian - Maastrichtian Lokoja Formation
- Pre-Cambrian – Lower Palaeozoic Basement Complex
- Studied Outcrop Locations

**Fig. 1**
Fig. 2
### POLLEN AND SPORES

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Fig. 5

- Fresh Water swamp
- Shallow marine
OCCURRENCE OF SOME MASTRICHTIAN DINOFLAGELLATE CYSTS FROM THE UPPER

Fig. 6

Shallow marine

Fresh water swamp

Thickness

Lithology

Sample No

PMI

FACIES INTERPRETATION

0 100 200 300 400
Fig. 7

Frequency chart for the marine microplanktons group

- Peridinoids
- Spiniferites
- Microforaminiferal linings
- Algae

Series 1
CONCLUSIONS

The palynomorph assemblage from the sections studied indicates predominance of terrestrial derived pollen and spores and significant representation of diverse marine dinoflagellates in the lower part of the section. The present data provide the first microplankton evidence of shallow marine condition in parts of the Patti Formation.

The sections of the Patti Formation is assigned Maastrichtian age and they correlate with part of Nkporo Shale in Anambra Basin and Gombe Formation in Gongola Basin, northeast Nigeria.

AKNOWLEDGEMENTS

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REFERENCES


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LIST OF FIGURES AND PLATE

Fig. 1 (a) Map of Nigeria showing NW/SE trending Bida Basin (b) Geological map of study area after Agyingi (1991) showing locations of studied sections.

Fig. 2 Regional stratigraphic successions in the Bida Basin and restored NW-SE-S stratigraphic relationships from the Bida basin to the Anambra Basin (Modified after Akande et al. 2005)

Fig. 3 Lithologic section and distribution of pollen and spore and dinoflagellates in the Patti Formation exposed at Ahoko, along Lokoja – Abuja highway.

Fig. 4 Lithologic sections and distribution of pollen and spores and dinoflagellates in the Patti formation exposed at Ahoko quarry.

Fig. 5 Palynological Marine Index (PMI) plot and facies interpretation of Patti Formation samples at Ahoko, along Lokoja – Abuja highway.

Fig. 6 Palynological Marine Index (PMI) plot and facies interpretation of Patti Formation exposed at Ahoko quarry.

Fig. 7 Frequency chart of the dinoflagellate groups in the study area (a) Pie chart (b) Bar chart. Plate 1

Fig. 1 Andalusiella polymorpha
Fig. 2 Senegalinium bicavatum
Fig. 3 Senegalinium bicavatum
Fig. 4 Senegalinium psilatum
Fig. 5 Deflandrea sp.
Fig. 6 Dinogymnium euclaensis
Fig. 7 Dinogymnium acuminatum
Fig. 8 Dinogymnium digitus
Fig. 9 Paleocystodinium australinum
Fig. 10 Microforaminifera lining
Fig. 11 Isabelidium cooksoni

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Fig. 13 Aeroligera senonensis
Fig. 14 Spiniferites ramosus
Fig. 15 Pediastrum
Fig. 16 Cristaeturites cristatus
Fig. 17 Buttinia andreevi
Fig. 18 Retidiporites magdalenensis

Magnifications; Figs 1,3,4,5,9,10,16 x 800, others x 1000.