



Provenance and Tectonic Setting of Leuma Field Sediments, Coastal Swamp Depobelt, Niger Delta Basin, Nigeria

*IGBINIGIE, NS; OGBAMIKHUMI, A

Department of Geology, University of Benin, Benin City, Nigeria.

**Corresponding Author Email: nosa.igbinigie@uniben.edu*

ABSTRACT: Geochemical analysis of sediments recovered from NS-1 well and NS-2 well, Leuma Field, Coastal Swamp Depobelt, Niger Delta Basin was done to establish tectonic setting and provenance. Data for ten major elements and forty-three trace elements were obtained from two boreholes. The geochemical signals display systematic stratigraphic trends in the two wells that depict one source terrain. To determine the provenance of NS-1 well and NS-2 well sediments, Th/Co vs La/Sc plot was utilized which inferred that the sediments from NS-1 and NS-2 wells were derived from felsic source rocks. TiO₂ versus Ni bivariate plot was also used to establish the provenance and it revealed that the source of the sediments penetrated by NS-1 and NS-2 wells is predominantly acidic in nature. The provenance of NS-1 and NS-2 wells sediments was further confirmed by considering the ratios of Thorium/Scandium (Th/Sc), Thorium/Cobalt (Th/Co), Chromium/Thorium (Cr/Th) and Lanthanum/Scandium (La/Sc). For NS-1 well, Thorium/Scandium (Th/Sc) range from 1.12-2.01, Thorium/Cobalt (Th/Co) range from 0.91-1.66, Chromium/Thorium (Cr/Th) range from 4.21-16.04 and Lanthanum/Scandium (La/Sc) range from 3.69-8.78. For NS-2 well, Thorium/Scandium (Th/Sc) range from 0.95-2.05, Thorium/Cobalt (Th/Co) range from 0.94-1.91, Chromium/Thorium (Cr/Th) range from 4.32-15.43 and Lanthanum/Scandium (La/Sc) range from 2.58-6.66. These values inferred that the sediments recovered from NS-1 and NS-2 wells were transported from felsic source rocks. Inorganic geochemical results infer that the tectonic setting for NS-1 and NS-2 wells facies is passive continental margin.

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The conventional objective of provenance studies is to reconstruct and interpret the history of sediment supply from initial erosion of a parent rock to the final burial of its detritus. Provenance study is done to deduce the geographic location and characteristics of the source area. Important factors such as the location and nature of source area, hinterland drainage pattern and pathways through which sediments have been transferred from source to basin influence the composition of the sedimentary rocks. This evolution may be recorded in the characteristics of the sediment that are deposited in a basin (Cox *et al.*, 1995; Nesbitt, 1990; Nesbitt and Young, 1982). Rare earth elements have historically been applied to many rock types to help decipher the origin and provenance evolution of rocks (Bhatia, 1985; Davies and Pickering, 1999; McLennan *et al.*, 1990). Low contents of Cr imply a

felsic provenance, and high levels of Cr and Ni are essentially found in sediments derived from ultramafic rocks (Wronkiewicz and Condie, 1990; Cullers and Podkovyrov, 2000; Nagarajan *et al.*, 2007). Ratios such as La/Sc, Th/Sc, Th/Co, and Th/Cr are significantly different in felsic and basic rocks and may possibly allow constraints on the average provenance composition (Cullers, 1994, 1995). Using major elemental compositions, provenance can be discriminated into mafic, intermediate, felsic igneous rocks and Quartzose sedimentary rocks fields. The ratio of Al₂O₃/TiO₂ in shales has been suggested to be similar to that of the source rock as such it is used as an index of provenance (Hayashi *et al.*, 1997). They stated that values greater than twenty-one imply sediments from felsic origin. (Floyd and Leveridge, 1987) proposed a plot of La/Th against Hf to

discriminate provenance. (McLennan *et al.*, 1990) plotted Th versus Sc to infer provenance. (Gao *et al.*, 1995) used Co/Th versus La/Sc diagram for provenance discrimination. Redox-sensitive trace element concentrations or ratios are among the main extensively used indicators of redox conditions in modern and ancient sedimentary deposits (Calver and Pedersen, 1993; Jones and Manning, 1994; Crusius *et al.*, 1996; Dean *et al.*, 1997, 1999; Yarincik *et al.*, 2000; Morford *et al.*, 2001; Pailler *et al.*, 2002; Algeo and Maynard, 2004). Plate tectonic processes impart a distinctive geochemical signature to sediments in two separate ways. Firstly, different tectonic environments have distinctive provenance characteristics and, secondly, they are characterized by distinctive sedimentary processes. Sedimentary basins may be assigned to the following tectonic settings for active continental margin, passive continental margin, oceanic island-arc, continental island-arc, and collisional setting. (Bhatia, 1985; Roser and Korsch, 1986) stated that the chemical compositions of clastic rocks are significantly controlled by plate tectonic

settings of their provenances, consequently clastic rocks from different tectonic settings possess terrain-specific geochemical signatures.

Geologic Background: The Niger Delta Basin ranks among the world’s most prolific petroleum producing Tertiary Deltas (Selley, 1997). It occupies the Gulf of Guinea continental margin in equatorial West Africa between Latitude 3⁰ and 6⁰ N and Longitude 5⁰ and 8⁰E. The Niger Delta is framed on the northwest by a subsurface continuation of the West African Shield, the Benin Flank. The eastern edge of the basin coincides with the Calabar Flank to the south of the Oban Masif (Murat, 1972). Well sections through the Niger Delta generally display three vertical lithostratigraphic subdivisions: an upper delta top facies; a middle delta front lithofacies; and a lower pro-delta lithofacies. These lithostratigraphic units correspond respectively with the Benin Formation (Oligocene-Recent), Agbada Formation (Eocene-Recent) and Akata Formation (Paleocene-Recent) of (Short and Stauble, 1967) respectively.

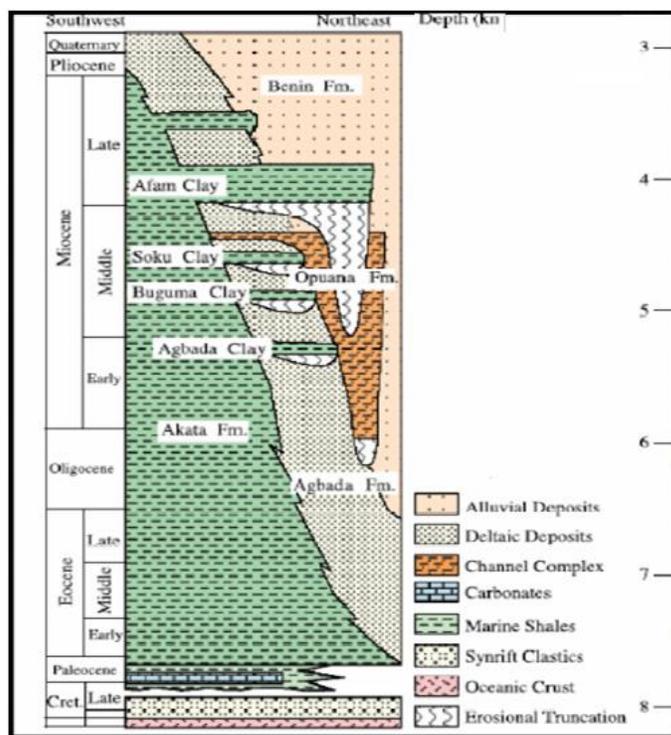


Fig 1: Stratigraphic column showing the three formations of the Niger Delta (modified after Doust and Omatsola, 1990).

The evolution of the Niger Delta was controlled by pre- and syn-sedimentary tectonics described by (Evamy *et al.*, 1978; Knox and Omatsola, 1989 and Stacher, 1995). The tectonic framework of the continental margin along the West Coast of equatorial Africa is controlled by Cretaceous fracture zones expressed as trenches and ridges in the deep Atlantic.

The fracture zone ridges subdivide the margin into individual basins, and, in Nigeria, form the boundary faults of the Cretaceous Benue-Abakaliki Trough, which cuts far into the West African shield. The trough represents a failed arm of a rift triple junction associated with the opening of the South Atlantic. Rifting started in the Late Jurassic and persisted into

the Middle Cretaceous (Lehner and De Ruiter, 1977). In the Niger Delta region, rifting diminished altogether in the Late Cretaceous and the gross paleogeography of the region as well as the relative position of the African and South American plates since rifting began. After rifting ceased, gravity tectonics became the primary deformational process. For any given depobelt, gravity tectonics were completed before deposition of the Benin Formation and are expressed in complex structures, including shale diapirs, roll-

over anticlines, collapsed growth fault crests, back-to-back features, and steeply dipping, closely spaced flank faults (Evamy *et al.*, 1978; Xiao and Suppe, 1992). These faults mostly offset different parts of the Agbada Formation and flatten into detachment planes near the top of the Akata Formation. The study wells (NS-1 well and NS-2 well) are located in Leuma Field, Coastal Swamp Depobelt, Niger Delta Basin. Figure 2 shows the location of NS-1 well and NS-2 wells.

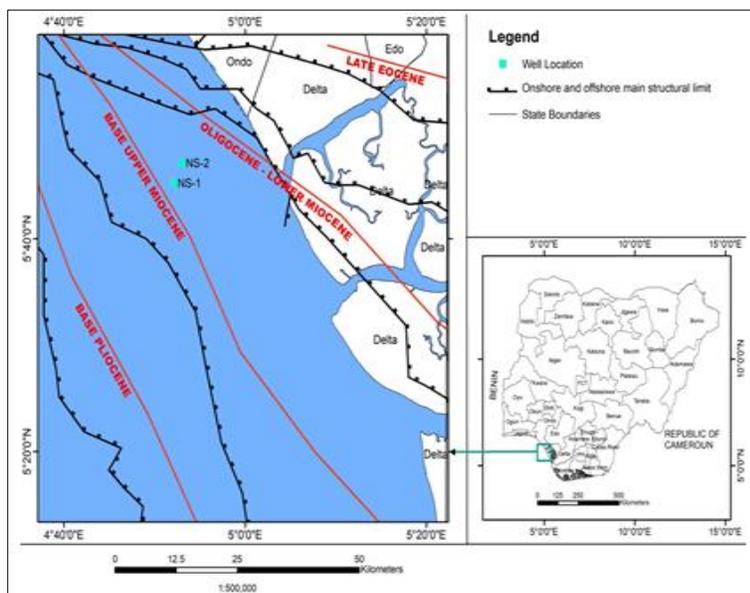


Fig 2: Map showing Location of NS-1 well and NS-2 well

Methodology: The pulverized ditch cutting samples were analyzed with X-ray fluorescence (XRF) and inductively coupled plasma mass spectrometry (ICP-MS) techniques. Twenty four samples (twelve from NS-1 well and twelve from NS-2 well) comprising of sands and shales were selected for this study using the aforementioned methods. The results derived from the analysis were used for the geochemical characterization of the wells. These analytical methods yielded results for ten (10) major elements, reported as oxide percent by weight (SiO_2 , Al_2O_3 , Fe_2O_3 , MgO , MnO , CaO , TiO_2 , Na_2O , K_2O and P_2O_5). Results for

forty-three (43) trace element (V, Co, Cr, Ni, Cu, Zn, Ga, Ge, As, Rb, Sc, Y, Zr, Nb, Mo, Lu, Ag, In, Sn, Sb, Cs, Ba, La, Ce, Pr, Nd, Sm, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb, Hf, Ta, W, Tl, Pb, Bi, Th, U) reported in ppm were also recorded. Although data for a total of 53 elements was acquired, tectonic setting and provenance study was done with few selected elements or ratios of elements (Pearce *et al.*, 2005a; Ratcliffe *et al.*, 2006). These key elements and element ratios, termed key indices, were used for inorganic geochemical characterization.

RESULTS AND DISCUSSION

Tables 1, 2, 3, and 4 show the findings of the investigation. Table 1 and table 2 display major element data for NS-1 and NS-2 wells, respectively, while table 3 and table 4 display trace element data for NS-1 and NS-2 wells. The findings were used to determine the sediments' provenance and tectonic setting.

Provenance Studies: Using major elemental compositions, provenance can be discriminated into mafic, intermediate, felsic igneous rocks and Quartzose sedimentary rocks fields as reported. As suggested by (Cullers, 2002), the plot of Th/Co versus La/Sc can be used to establish the provenance of sediments. To determine the provenance of NS-1 and NS-2 sediments, the Th/Co vs. La/Sc plot of Cullers,

2002) was utilized which inferred that the sediment from NS-1 and NS-2 wells were derived from felsic source rocks (Fig 3). Furthermore, the concept of (Floyd *et al.*, 1989) was also adopted to establish the source of the sediments penetrated by NS-1 well and NS-2 well. TiO₂ versus Ni bivariate plot was used and it revealed that the source of the sediments penetrated by NS-1 and NS-2 wells is predominantly acidic in nature (Fig 4). The provenance of NS-1 and NS-2 wells sediments was further confirmed by considering the ratios of Thorium/Scandium (Th/Sc), Thorium/Cobalt (Th/Co), Chromium/Thorium

(Cr/Th), Lanthanum/Scandium (La/Sc). The ratios of these elements were calculated from the result shown in table 5, For NS-1 well, Thorium/Scandium (Th/Sc) range from 1.12-2.01, Thorium/Cobalt (Th/Co) range from 0.91-1.66, Chromium/ Thorium (Cr/Th) range from 4.21-16.04 and Lanthanum/Scandium (La/Sc) range from 3.69-8.78. For NS-2 well, Thorium/Scandium (Th/Sc) range from 0.95-2.05, Thorium/Cobalt (Th/Co) range from 0.94-1.91, Chromium/Thorium (Th/Cr) range from 4.32-15.43 and Lanthanum/Scandium (La/Sc) range from 2.58-6.66.

Table 1: Major Elements of the Selected Sandstone and Shales from NS-1 Well

Sample Number	Depth Interval (ft)	SiO ₂ (%)	Al ₂ O ₃ (%)	Fe ₂ O ₃ (%)	MnO (%)	MgO (%)	CaO (%)	Na ₂ O (%)	K ₂ O (%)	TiO ₂ (%)	P ₂ O ₅ (%)
1(SHALE)	4920-4950	75.82	9.82	5.26	0.05	0.72	0.71	0.82	2.10	1.31	0.26
2(SHALE)	5220-5250	75.66	8.44	6.04	0.06	0.74	0.82	0.91	1.91	1.09	0.31
3(SHALE)	5280-5310	75.13	9.81	7.72	0.05	0.65	0.54	0.83	1.83	1.06	0.39
4(SAND)	5400-5430	89.76	2.56	2.95	0.02	0.14	0.15	0.81	0.29	0.17	0.03
5(SAND)	5460-5490	90.31	2.78	3.02	0.03	0.13	0.15	0.69	0.31	0.17	0.03
6(SHALE)	5700-5730	74.21	10.72	6.98	0.06	0.52	0.69	0.79	2.00	0.98	0.29
7(SHALE)	6180-6210	76.95	7.64	7.91	0.05	0.77	0.52	0.77	1.76	0.87	0.32
8(SAND)	6600-6630	91.44	2.30	3.01	0.02	0.15	0.12	0.73	0.34	0.19	0.04
9(SAND)	6660-6690	89.93	2.53	3.04	0.04	0.12	0.14	0.82	0.28	0.18	0.03
10(SAND)	7260-7290	91.64	2.36	2.88	0.02	0.11	0.16	0.76	0.36	0.17	0.03
11(SHALE)	7440-7470	73.41	8.63	7.24	0.05	0.87	0.66	0.74	1.65	1.04	0.91
12(SAND)	7500-7530	89.87	2.72	3.00	0.03	0.16	0.17	0.64	0.29	0.16	0.03

Table 2: Major Elements of the Selected Sandstone and Mudstone from NS-2 Well

Sample Number	Depth Interval (ft)	SiO ₂ (%)	Al ₂ O ₃ (%)	Fe ₂ O ₃ (%)	MnO (%)	MgO (%)	CaO (%)	Na ₂ O (%)	K ₂ O (%)	TiO ₂ (%)	P ₂ O ₅ (%)
1(SHALE)	5010-5040	75.6	9.36	5.67	0.05	0.67	0.77	0.17	0.88	1.37	0.04
2(SHALE)	5250-5280	76.7	10.14	6.78	0.06	0.56	0.82	0.21	0.71	0.97	0.02
3(SAND)	5490-5520	87.7	2.72	2.86	0.02	0.21	0.19	0.19	0.92	0.16	0.16
4(SAND)	5820-5850	85.2	2.65	3.01	0.03	0.35	0.18	0.18	0.86	0.17	0.18
5(SHALE)	6120-6150	74.4	8.75	4.98	0.05	0.21	0.61	0.22	0.96	0.86	0.19
6(SAND)	6300-6330	90.1	2.86	2.75	0.04	0.28	0.19	0.27	0.77	0.18	0.07
7(SAND)	6370-6600	87.8	2.55	3.66	0.04	0.36	0.16	0.17	0.87	0.18	0.05
8(SHALE)	6750-6780	75.5	9.24	5.22	0.06	0.55	0.56	0.28	0.79	1.46	0.14
9(SHALE)	7260-7290	74.7	9.94	2.76	0.05	0.37	0.78	0.17	0.66	1.13	0.14
10(SAND)	7350-7380	90.4	2.93	3.44	0.02	0.19	0.15	0.17	0.97	0.15	0.02
11(SAND)	7620-7650	89.3	2.47	5.37	0.03	0.34	0.16	0.27	0.98	0.16	0.03
12(SHALE)	7860-7890	75.1	8.56	5.37	0.06	0.48	0.67	0.16	1.32	0.96	0.16

Table 3: Results of selected trace elements for NS-1 well

Sample Number	Depth Interval (Ft)	Co	Cr	Ni	Sc	Th	La
1(SHALE)	4920-4950	14.7	82	35	12	13.4	56.7
2(SHALE)	5220-5250	14.3	87	41	14	15.7	59.2
3(SHALE)	5280-5310	13.8	93	35	11	18.9	70.3
4(SAND)	5400-5430	3.8	35	20	2.1	4.12	10.2
5(SAND)	5460-5490	3.9	40	18	1.9	3.81	12.1
6(SHALE)	5700-5730	12.1	99	45	17	20.1	62.7
7(SHALE)	6180-6210	11.9	87	37	16	18.2	81.2
8(SAND)	6600-6630	4.0	49	10	2.2	3.66	12.9
9(SAND)	6660-6690	3.6	52	15	2.0	3.85	12.3
10(SAND)	7260-7290	3.2	60	14	2.1	3.74	114.7
11(SHALE)	7440-7470	12.3	83	36	14	19.7	83.5
12(SAND)	7500-7530	3.1	57	11	1.8	3.56	15.8

Table 4: Results of selected trace elements for NS-2 well

Sample Number	Depth Interval (Ft)	Co	Cr	Ni	Sc	Th	La
1(SHALE)	5010-5040	14.9	85	37	15	14.3	56.3
2(SHALE)	5250-5280	13.3	80	43	12	15.4	61.5
3(SAND)	5490-5520	3.8	37	20	2.5	4.45	11.3
4(SAND)	5820-5850	3.9	45	18	2.0	3.95	13.4
5(SHALE)	6120-6150	12.8	97	34	13	17.8	73.7
6(SAND)	6300-6330	4.0	48	10	2.6	3.77	12.8
7(SAND)	6570-6600	3.6	52	15	2.0	3.87	12.3
8(SHALE)	6750-6780	11.1	92	42	18	21.3	67.8
9(SHALE)	7260-7290	10.5	87	37	17	19.7	85.1
10(SAND)	7350-7380	3.2	59	14	1.9	3.91	14.1
11(SAND)	7620-7650	3.1	54	11	2.7	3.50	15.9
12(SHALE)	7860-7890	12.6	95	39	13	20.3	86.7

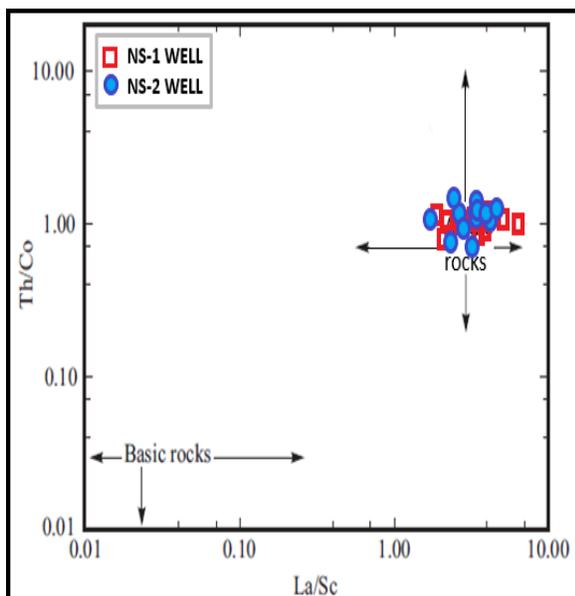


Fig 3: Th/Co versus La/Sc diagram for NS-1 and NS-2 wells. (After Culler, 2002)

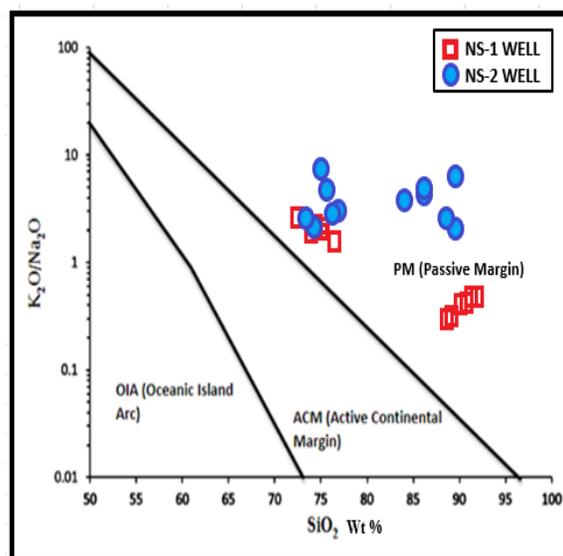


Fig 5: Tectonic discrimination plot for NS-1 and NS-2 wells. After Roser and Korsch (1986). PCM: passive continental margin, ACM: active continental margin and OIA: oceanic island arc.

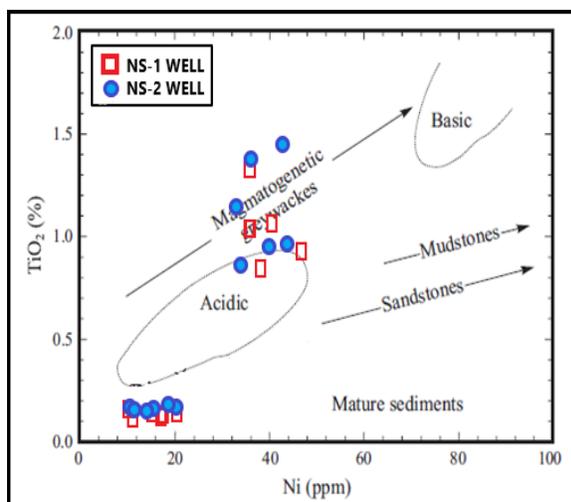


Fig 4: TiO₂ versus Ni bivariate of NS-1 and NS-2 wells (After Floyd *et al.* 1989)

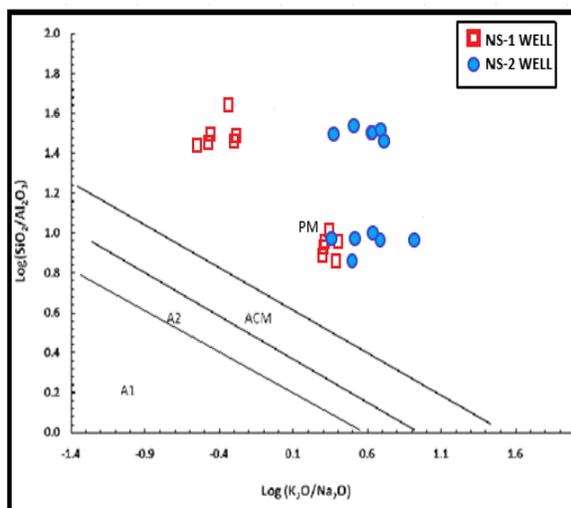


Fig 6: Log (K₂O/Na₂O) ratio versus Log (SiO₂/Al₂O₃) ratio tectonic discrimination diagram of NS-1 and NS-2 wells. After Maynard *et al.* (1982). A1 - arc setting and andesitic detritus, A2 - evolved arc setting, felsic pluton detritus, ACM - Active continental margin, PM - Passive Continental Margin.

As prescribed by Taylor and McLennan (1985), Cullers (1994), (2000), the values recorded above were compared to those of sediments resulting from typical felsic and basic rocks as well as to upper

continental crust (UCC) and PAAS values, it was inferred that the sediments recovered from NS-1 and NS-2 wells were transported from felsic source rocks.

Table 5: Range of Elemental Ratios of NS-1 compared to the Ratios in Similar Fractions derived from Felsic and Mafic Rocks, Upper Continental Crust (UCC) and Post-Archean Australian Average Shale. After Cullers (1994) (2000), Taylor and McLenna (1985).

Elemental Ratio	NS-1 Well Range	Range For Felsic Rocks	Range For Mafic Rocks	PAAS	UCC
Th/Sc	1.12-2.01	0.84-20.5	0.05-0.22	0.90	0.79
Th/Co	0.91-1.66	0.67-19.4	0.04-1.40	0.63	0.63
Cr/Th	4.21-16.04	4.0-15.0	25-100	7.53	7.76
La/Sc	3.69-8.78	1.2-6.6	0.43-0.86	2.40	2.21

Table 6: Range of Elemental Ratios of NS-2 well compared to the ratios in similar fractions derived from felsic and mafic rocks, Upper Continental Crust (UCC) and Post-Archean Australian Average Shale. After Cullers (1994) (2000), Taylor and McLenna (1985).

Elemental Ratio	NS-2 Well Range	Range For Felsic Rocks	Range For Mafic Rocks	PAAS	UCC
Th/Sc	0.95-2.05	0.84-20.5	0.05-0.22	0.90	0.79
Th/Co	0.94-1.91	0.67-19.4	0.04-1.40	0.63	0.63
Cr/Th	4.32-15.43	4.0-15.0	25-100	7.53	7.76
La/Sc	2.58-6.66	1.2-6.6	0.43-0.86	2.40	2.21

Tectonic Setting: The concept of (Roser and Korsch, 1986) was applied to determine the tectonic setting of NS-1 and NS-2 wells sediments. (Roser and Korsch, 1986) plotted K_2O/Na_2O vs SiO_2 to determine the provenance of sediments. The recognized tectonic settings on the K_2O/Na_2O versus SiO_2 discrimination diagram of (Roser and Korsch, 1986) are: the passive continental margin (PCM), active continental margin (ACM) and oceanic island arc (OIA). When applied for the samples recovered from NS-1 and NS-2 wells, they plotted mainly in the passive continental margin zone. It was therefore inferred that the tectonic setting for NS-1 and NS-2 well facies is passive continental margin (Figure 5) Log (K_2O/Na_2O) versus Log (SiO_2/Al_2O_3) was used to determine the tectonic setting as proposed by (Maynard *et al.*, 1982). The recognized tectonic settings on the Log (K_2O/Na_2O) ratio versus Log (SiO_2/Al_2O_3) ratio discrimination diagram of (Maynard *et al.*, 1982) are: A1 - arc setting and andesitic detritus; A2 - evolved arc setting, felsic pluton detritus; ACM - Active continental margin; PM - passive margin. When utilized for the samples of NS-1 and NS-2 wells, they plotted mainly in the passive margin zone which infer that the tectonic setting for the NS-1 and NS-2 wells facies is passive continental margin (Figure 6).

Conclusion: The provenance and tectonic setting of the sandstones and shales from NS-1 well and NS-2 well, Leuma Field, Coastal Swamp Depobelt, Niger Delta Basin were investigated with geochemical methods. The study revealed that the sediments from the two wells were derived from felsic/acidic source. Tectonic studies reveal that the sediments were deposited on a passive margin setting that received most of its detritus from the nearby Southwestern

Nigeria Basement Complex Rocks. From the tectonic history, it is believed that the studied field was not affected by major tectonic activities.

REFERENCES

- Algeo, TJ; Maynard, JB (2004). Trace-element behavior and redox facies in core shales of Upper Pennsylvanian Kansas-type cyclothems. *Chem Geol* 206: 289-318.
- Bhatia, MR (1985). Rare earth element geochemistry of Australian Paleozoic graywackes and mudrocks: Provenance and Tectonic control: *Sedimentary geology*. 45: 97-113.
- Calvert, SE; Pedersen, TF (1993). Geochemistry of recent oxic and anoxic marine sediments: implications for the geological record. *Mar Geol* 113: 67-88.
- Cox, R; Lowe, DR; Cullers, RL (1995). The influence of sediment recycling and basement composition on evolution of mudrock chemistry in the southwestern United States. *Geochim Cosmochim Acta* 59: 2919-2940.
- Crusius, J; Calvert, S; Pedersen, T; Sage, D (1996). Rhenium and molybdenum enrichments in sediments as indicators of oxic, suboxic and sulfidic conditions of deposition. *Earth Planet Sc* 145: 65-78.
- Cullers, RL (1995). The controls on the major and trace element evolution of shales, siltstones and sandstones of Ordovician to Tertiary age in the

- Wet Mountain region, Colorado, U.S.A. *Chemical Geology*, 123(1-4): 107-131.
- Cullers, RL (1994). The controls on the major and trace element variation of shales, siltstones and sandstones of Pennsylvanian–Permian age from uplifted continental blocks in Colorado to platform sediment in Kansas, USA. *Geochimica et Cosmochimica Acta*, 58(22): 4955-4972.
- Cullers, RL; Podkovyrov, VN (2000). Geochemistry of the Mesoproterozoic Lakhanda shales in southeastern Yakutia, Russia: implications for mineralogical and provenance control, and recycling. *Precambrian Res* 104: 77-93.
- Dean, WE; Gardner JV; Piper, DZ (1997). Inorganic geochemical indicators of glacial-interglacial changes in productivity and anoxia of the California continental margin. *Geochim Cosmochim Acta* 61: 4507-4518.
- Doust, B; Omatsola, E (1990). Niger Delta, in, Edwards, J. D., and Santogrossi, P.A., eds., *Divergent/passive Margin Basins.*, AAPG *Memoir 48: Tulsa, American Association of Petroleum Geologists.* 239-248.
- Evamy, BD; Haremboure, J; Kamerling, P; Knaap, A; Malloy, F; Rowlands, PH (1978). Hydrocarbon habitat of Tertiary Niger Delta. *AAPG Bulletin*, 62: 1-39.
- Floyd, PA; Leveridge, BE (1987). Tectonic environment of the Devonian Gramscatho basin, south Cornwall: framework mode and geochemical evidence from turbiditic sandstones. *Journal of the Geological Society*. 144(4): 531-542.
- Gao, S; Zhang, BR; Gu, XM; Xie, QL; Gao, CL; Guo, XM (1995). Silurian-Devonian provenance changes of South Qinling basins: implications for accretion of the Yangtze (South China) to the North China cratons. *Tectonophysics*, 250(1-3): 183-197.
- Hayashi, KI; Fujisawa, H; Holland, HD; Ohmoto, H (1997). Geochemistry of ~ 1.9 Ga sedimentary rocks from northeastern Labrador, Canada. *Geochimica et cosmochimica acta*, 61(19): 4115-4137.
- Jones, B; Manning, DAC (1994). Comparison of geological indices used for the interpretation of palaeoredox conditions in ancient mudstones. *Chem Geol* 111: 111-129.
- Knox, GJ; Omatsola, EM (1989). Development of the Cenozoic Niger Delta in terms of the 'Escalator Regression' model and impact on hydrocarbon distribution. In: van der Linden, W.J.M., Cloetingh, S.A.P.L., Kaasschieter, J.P.K., van der Graff, W.J.E., Vandenberghe, J., van der Gun, J.A.M. (Eds.), *Proceedings KNGMG Symposiums Coastal Lowlands, Geology and Geotechnology. Kluwer Academic Publishers, Amsterdam*, 181-202.
- Lehner, P; De Ruiter, PAC (1977) "Structural History of Atlantic Margins of Africa", *AAPG Bulletin*, 6 (7): 961-981
- Morford, JL; Russell, AD; Emerson, S (2001). Trace metal evidence for changes in the redox environment associated with the transition from terrigenous clay to diatomaceous sediment, Saanlich Inlet, BC. *Mar Geol* 174: 355-369.
- Morley, R.J. and Richards, K. (1993). Graminae cuticle: a key indicator of Late Cenozoic climatic change in the Niger Delta. *Review of Palaeobotany and Palynology* 77: 119-127
- Mclennan, SM; Hemming, S; McDaniel, DK; Hanson, GN (1990). Geochemical approaches to sedimentation, provenance, and tectonics. Processes controlling the composition of clastic sediments (Johnson, M.J. and Basu, A., eds.), 21-40, *Geological Society of America special paper* 284.
- Murat, RC (1972). Stratigraphy and Paleogeography of the Cretaceous and Lower Tertiary in Southern Nigeria. In T. F. Dessauvage, & A. J. Whiteman (Eds.), *African Geology*. 251-266.
- Nagarajan, R; Madhavaraju, J; Nagendral, R; Armstrong-Altrin, JS, Moutte, J (2007). Geochemistry of Neoproterozoic shales of the Rabanpalli Formation, Bhima Basin, Northern Karnataka, southern India: implications for provenance and palaeoredox conditions. *Rev Mex Cienc Geol* 24: 150-160.
- Nesbitt, HW; Young, GM (1982). Early Proterozoic Climate and Plate motion inferred from major element chemistry of lutites, *Nature* 299: 715 – 717.

- Pailler, D; Bard E; Rostek, F; Zheng, Y; Mortlock, R; Van-Geen A (2002). Burial of redox-sensitive metals and organic matter in the equatorial Indian Ocean linked to precession. *Geochim Cosmochim Acta* 66: 849-865.
- Pearce, TJ; Wray, DS; Ratcliffe, KT; Wright, DK; Moscariello, A (2005a). Chemostratigraphy of the Upper Carboniferous Schooner Formation, southern North Sea. In: Carboniferous hydrocarbon geology: the southern North Sea and surrounding onshore areas. In: Collinson, J.D., Evans, D.J., Holliday, D.W. and Jones N.S. (eds) Yorkshire Geological Society, *Occasional Publications series*, 7, 147–64
- Ratcliffe, KT; Wright, AM; Hallsworth, C; Morton, A; Zaitlin, BA; Potoki, D; Wray, DS (2004). Alternative correlation techniques in the petroleum industry: an example from the (Lower Cretaceous) Basal Quartz, Southern Alberta. *Bulletin of the American Association of Petroleum Geologists*, 88, 1,419–32.
- Roser, BP; Korsch, RJ (1986). Determination of tectonic setting of sandstone-mudstone suites using SiO₂ content and K₂O/Na₂O ratio. *The J. Geol.* 94(5), 635-650.
- Stacher, P. (1995). Present understanding of the Niger Delta hydrocarbon habitat, in: Oti, MN; Postma, G, eds. *Geology of Deltas: Rotterdam, AA. Balkema*, 257-267.
- Selley, R. (1997). Sedimentary basins of the world. In K. Hsu (Ed.), *The basins of northwest Africa: structural; evolution* (pp. 17-26). Elsevier, Amsterdam.
- Short, KC; Stauble, A (1967). Outline of the geology of the Niger Delta. *AAPG Bulletin*, 51, 761-779
- Wronkiewicz, DJ; Condie, KC (1990). Geochemistry and mineralogy of sediments from the Ventersdorp and Transvaal Supergroups, South Africa: Cratonic evolution during the early Proterozoic. *Geochim Cosmochim Acta* 54: 343-354.
- Xiao, H; Suppe, J (1992) " and Origin of Roll Over" *AAPG Bulletin*. 96 (4): 509-529
- Yarincik, KM; Murray, RW; Lyons, TW; Peterson, LC; Haug, GH (2000). Oxygenation history of bottom waters in the Cariaco Basin, Venezuela, over the past 578,000 years: results from redoxsensitive metals (Mo, V, Mn, and Fe). *Paleoceanography* 15:593-604.