Advances in optical earth observation for geological mapping: A Review

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Abstract

Recent increases in the availability of earth observation (EO) data and the advances made in its processing have opened up new opportunities for earth monitoring studies. This article reviews recent advances in EO, specifically in the optical domain, and comments on their potential for geological investigations. A review of the literature revealed that much work has been done on the use of multispectral to hyperspectral EO data for geological investigations, but that relatively little emphasis has been placed on the value of high spatial resolution in geological studies. From the literature it was found that geographic object based image analysis (GEOBIA) is a relatively new paradigm in remote sensing that has been shown to reduce the intra-class spectral variability in high spatial resolution imagery, thereby increasing image classification accuracies. It has been used for many applications, but relatively little research has been done on the application of GEOBIA for geological investigations. Another finding of this review is that the power of EO data increases substantially when different types of data are used (fused) together. Data fusion seizes the advantages of the capabilities of each sensor while reducing its limitations. By its nature, GEOBIA offers the ability to fuse disparate data sources and the review concludes that more research is required on the use of GEOBIA and data fusion of high resolution EO data for geological mapping.

1. Introduction

Remote sensing (RS) of the earth, or earth observation (EO), can be defined as the science of the identification of earth-surface features and the estimation of their geophysical and biophysical properties using electromagnetic radiation, measured by sensors on airborne (aircraft) or spaceborne (satellites, space shuttles) platforms (Agar and Coulter, 2007; Drury, 2001). EO has played an important role in geological mapping, especially in areas with limited infrastructure and harsh environmental conditions. EO methods offer: the opportunity to analyse and map surface geology in a relatively short time and at reduced cost; provide additional (sometimes new) information for preliminary geological investigations; and give a synoptic view of a study area often difficult to obtain from field-based observation alone (Chernicoff and Nash, 2002; Kariuki *et al.*, 2004; Salati *et al.*, 2011). Broad lithological information is deduced

from a variety of indicators observed in remotely sensed imagery, including landforms, structural features, soils, vegetation types and drainage patterns (Agar and Coulter, 2007; Gad and Kusky, 2006; 2007; Hewson et al., 2005; Rajesh, 2004). Many authors have studied the strong correlation between lineament interpretation and analysis of remotely sensed imagery, leading to discoveries of major mineral deposits of great economic value (Fernades da Silva et al., 2005; Hashim et al., 2013; Meijerink, 1996; Rajesh, 2004; Rinaldi, 2007; Saadi et al., 2011). Mineral exploration often employs EO to map fault fractures because of their known geological and spatial association with ore deposits, together with frequent targeting of indicator minerals, which often highlight the alteration mechanisms of economic mineral deposits (Azizi et al., 2010; Gabr et al., 2010; Pour and Hashim 2011a; 2011b; Sabins, 1999). A number of useful reviews of EO applications are available (Agar and Coulter, 2007; Bhan and Krishnanunni, 1983; Goetz, 2009; Koirala, 1997; Lipton 1997; Navalgund et al., 2007; Rajesh, 2004; Sabins, 1999; Van der Meer et al., 2012). This paper expands on the existing reviews by considering recent developments in RS which could potentially aid geological investigations. The focus will be on improvements in spatial and spectral resolution of optical imagery and the value of data fusion for geological applications.

2. Overview of EO techniques

The application of EO has increased tremendously since its development in the early nineteenth century. Its capability to measure electromagnetic energy in the visible short-wave infrared (VSIR), thermal infrared (TIR) and microwave wavelength bands has led to the application of EO to geological investigations. Earth surface materials can be remotely detected by studying the way in which they reflect wavelengths with different intensities (Drury, 2001; Goetz, 2009; Sabins, 1999).

An EO sensor's ability to record the reflected intensities is based on its spectral and spatial resolution features. The spectral resolution of a sensor is defined by the number of spectral bands – ranging from a few (multispectral) to hundreds (hyperspectral) – a sensor can record. Spatial resolution denotes the smallest area of the ground surface for which electromagnetic energy can be recorded. Figure 1 provides an overview of some of the prominent EO sensors and their respective spectral ranges used for various geological investigations and other earth monitoring studies.

				Remote Sensing Wavelength for Geological mapping
				Visible short wave region (VSWIR) V-VISIBLE NRT-NEAT SWIR-SHORT-WAVE INFRARED TIR-THERMAL INFRARED 0.4 0.7 1.1 1.4 1.7 2.0 2.3 2.63 6 9 12 15.10 ³
				Wavelength in micrometres
PLATFORM		SENSOR	SPATIAL RESOLUTION	SPECTRAL RESOLUTION
PHOTOGRAPH	AIR- CRAFT	AERIAL PHOTOGRAPHS	Variable	Film development
MULTISPECTRAL	SATELLITE (ST)	LANDSAT MULTISPECTRAL SCANNER (MSS), 1-4	80 m VNIR	1234
		LANDSAT THEMATIC MAPPER (TM), 5	30 m VSWIR 120 m TIR	129 4 5 7 6
		LANDSAT THEMATIC MAPPER (ETM +), 7	15 m Pan 30 m VSWIR 120m TIR	PAN 122 4 5 7 6
		LANDSAT THEMATIC MAPPER (ETM +), 8 OLI & TIRS	15 m Pan 30 m VSWIR	PAN 1234 5 9 6 7 1011
		ASTER	15 m VNIR 30 m SWIR 90 m TIR	1 2 3 4 5 111 9 10 114
		SPOT	2.5 m 5 m 10 m Pan 20 mVNIR	PAN 12 3
		IKONOS	1 m Pan 4 m VNIR	PAN 1234
		WORLDVIEW 3	1,38 m VSWIR 4,10 m SWIR	12345678941114 56
HYPERSPECTRAL	rz	HYPERION	30 m VSWIR	220 BANDS
	AIRCRAFT	НҮМАР	5 m VSWIR	126 BANDS
		AVIRIS	20 m VSWIR	224 BANDS
		SEBASS	5 m TIR	128 128

Figure 1. Optical earth observation sensors with their spectral and spatial resolution characteristics (adapted from Mauger, 2003)

The required spectral and spatial resolutions depend on the application. A high spatial resolution, for instance, is often favoured in urban monitoring applications, while a high spectral resolution is often of more value in geological applications (Drury, 2001; Goetz, 2009; Sabins, 1999; Van der Meer, *et al.*, 2012). High spectral resolution is needed in many geological applications since they rely on the detection of unique spectral signatures of different rock forming minerals (Figure 2), caused by electronic transitions and ion vibrations of such minerals (Drury, 2001; Kruse and Lefkoff, 1993; Navalgund *et al.*, 2007; Rajesh, 2004; Sabins, 1999). Electronic transitions are most often created by iron, while vibrational transitions are often created by water, hydroxyl ions and carbonates (Rowan *et al.*, 2004). The absorption feature indicating iron oxide electronic transitions is reflected by wavelengths in the range of 0.35 to 1.5μm, while charge-transfer transitions are apparent at wavelengths shorter than 0.55μm. Vibration transitions produce reflectance anomalies in the near-infrared region of the spectrum between 1.1μm and 2.5μm, while Mg-OH also produces distinctive absorption features in this range. Vibration transitions associated with carbonates yield absorption features at about 2.3μm

(Drury, 2001; Rajesh, 2004).

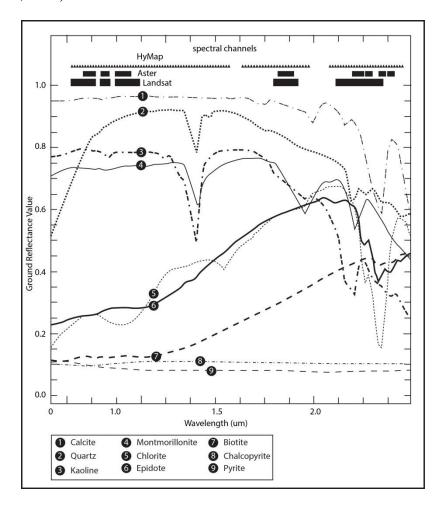


Figure 2. Mineral spectral signatures of some common rock forming minerals (adapted from Cocks *et al.*, 1998)

Various techniques for the analysis of EO imagery for geological investigations are available. Of these the most popular approaches are band ratios (Assiri *et al.*, 2008; Drury, 2001; Gabr *et al.*, 2010; Pour and Hashim, 2012a; 2012b); principal component analysis (PCA) (Gad and Kusky, 2007; Lipton, 1997; Rajesh, 2008; Sabins, 1999) and image classification (Brogaard and Prieler, 1998; Campbell, 1996; Gibson and Power, 2000).

The application of TIR EO data for geological investigations is popular owing to its ability to record the earth's surface emissivity (or temperature) and the rate of heat loss of different surface materials, which often indicates geophysical properties. TIR images are usually recorded in the range of 3 to 15μm, although the absorption by atmospheric gases restricts thermal sensing in the regions of 3 to 5μm and 8 to 14μm (Drury, 2001; Rowan and Mars, 2003; Van der Meer *et al.*, 2012; Vaughan *et al.*, 2003). According to Hecker (2006), TIR data complements the information provided by VSIR, particularly in geological studies where the distribution of

silicate minerals is important. The absorption features of silicate minerals are observed around 8 to 12µm wavelength thus making TIR systems invaluable for silicate mineral detection.

3. Advances in EO image processing

The recent advances in EO have concentrated on the development and processing of imaging spectrometry, also known as hyperspectral imaging. The development of hyperspectral imaging has been documented by several authors (Cloutis, 1996; Goetz, 2009; Kruse and Lefkoff, 1993; Plaza *et al.*, 2009). The acquisition capability of hundreds of narrow contiguous spectral bands enhances the identification of earth-surface materials by their characteristic reflectance or absorption discriminatory features. An example of such a sensor is the airborne visible infrared imaging spectrometer (AVIRIS), developed in 1984 with 224 contiguous spectral bands from 4.0 to 2.5μm at 20m spatial resolution. The HyMap hyperspectral imaging sensor, developed by the Australian company HyVista (Figure 1), is another example. The sensor covers the 0.45 to 2.5μm wavelength region in 126 spectral bands at 5m spatial resolution (Cocks *et al.*, 1998). Both the AVIRIS and HyMap operate on an aircraft platform. Currently, the only earthorbiting hyperspectral imaging sensor is Hyperion on the Earth Observation-1 (EO-1) satellite which was launched in the year 2000. It has the ability to acquire data in 220 spectral bands from 0.356 to 2.5μm at 30m spatial resolution (Kruse *et al.*, 2003).

The advent of high spectral resolution sensors led to the development of a suite of techniques for extracting spectral information, the goal being the derivation of surface compositional information of the earth's surface (Cloutis, 1996; Van der Meer *et al.*, 2012). Two broad categories of analytical techniques were developed and widely used for classifying hyperspectral data, namely spectrum-matching techniques and subpixel unmixing techniques. Spectrum-matching techniques express the spectral similarity of reference (derived from spectral library or field spectra of known materials) to test (image) spectra. This technique is applied through the spectral angle mapper (SAM) which measures spectral similarities by calculating the angle between the reference spectra and the test spectra (Bertels *et al.*, 2005; Kruse 1998; Van der Meer, 2006). The subpixel unmixing techniques aim to quantify the relative abundance (fractions, percent or area) of various materials within a pixel. This is based on the premise that a pixel may represent multiple surface materials, resulting in mixed (averaged) spectral signatures. Subpixel unmixing techniques separate such average spectral signatures into pure spectral

signatures which represent each material covered by the pixel.

The spatial resolution is also vital as it determines which earth-surface objects are detectable (Yu *et al.*, 2006). Finer (or higher) spatial resolution images provide greater spatial detail and therefore enable the delineation of small features compared to coarse (or low) spatial resolution images (Aplin, 2006; Mather, 2003; Woodcock and Strahler, 1987). Consequently, the spatial resolution of an image directly influences the scale at which it can be applied. The gradual refinement of geological work from scales of 1:250 000 to 1:50 000, and even 1:10 000 to 1:5 000, requires EO images with increasing spatial resolutions. For instance, early applications of EO to geological applications made use of Landsat (MSS) imagery with 80m resolution (Agar and Coulter, 2007; Bhan and Krishnanunni, 1983). The availability of 30m resolution Landsat (TM) imagery has enabled larger scale mapping (Gad and Kusky, 2006; Mohammad *et al.*, 2001), while 5m resolution HyMap data is suitable for even larger scale surveys (Cloutis, 1996; Cocks *et al.*, 1998; Yu *et al.*, 2006).

Current advances in EO are providing images with increasing spatial resolutions of up to 31cm (World View 3 panchromatic band, launched on the 13th of August 2014). However, the inherent problem with the use of high spatial resolution imagery is its increase in intraclass spectral variability within homogenous earth-surface features (e.g. land-cover types) which decreases the statistical separability of such features in spectral data space and subsequently reduces the accuracy of image classifications (Aplin et al., 1997; Carleer et al., 2005; Woodcock and Strahler, 1987, Yu et al., 2006). To overcome the intraspectral variability caused by improved spatial resolution, geographic (also known as geospatial) object-based image analysis (GEOBIA) has been proposed (Blaschke, 2010; Blaschke and Hay, 2001; Hay and Castilla, 2008). GEOBIA involves partitioning a scene into discrete entities or segments from which meaningful image objects, based on their spatial and spectral attributes, can be generated (Desclee et al., 2006; Kim et al., 2009). The basic assumption of GEOBIA is that the image being classified is made up of relatively homogenous patches that are larger than individual pixels. To analyse EO images, GEOBIA starts by grouping pixels into image-objects (Hay and Castilla, 2008; Chen et al., 2008; Darwish et al., 2003; Orkonselenge, 2004) called image segmentation, a fundamental step in GEOBIA.

Image segmentation diminishes the within-object variations, while emphasizing the inter-object variations (Benz et al., 2004; Blaschke, 2010). This is achieved by the

merging of pixels into objects and is based on a local homogeneity criterion which defines similarities between adjacent pixels following a combination of colour (spectral value) and spatial properties (size, texture and shape). Various segmentation algorithms are available in literature, with Mishra *et al.* (2011) categorizing them (segmentation) into four types, namely edge-based, cluster-based, region-based and split-merge approaches. In *edge-based* approaches an image is partitioned based on abrupt changes in intensity. Image edges are detected and linked into contours that represent boundaries of image objects. In *cluster-based* approaches image segmentation stores image pixels in ascending order of intensity. A predefined number of clusters are separated by using an intensity histogram. In *region-based* approaches, the segments with similar intensities are merged according to a decision rule, while *split/merge* approaches are based on a colour and texture algorithm.

4. Examples of EO in geological applications

Notable publications in which EO was used for geological applications are summarized in this review. Chernicoff and Nash (2002) combined a 1-4-7 composite of Landsat Thematic Mapper (TM) bands with aeromagnetic data for analysing geological structures in the northern Precordillera region of Argentina. The information derived from the Landsat (TM) imagery conformed to published maps and also revealed a new set of previously unknown structures. The resulting structures were superimposed on an aeromagnetic image, which enabled the precise delineation of lithomagnetic domain boundaries and magnetic pattern breaks (inferred faults). The combination of these datasets led to the definition of a major structural system which controls the economic potential of gold targets in the study area. Hewson et al. (2005) used ASTER data to generate new, accurate and seamless geological and mineralogical information in Australia. Major minerals rich in Al-OH, Mg-OH, carbonates and quartz abundances were targeted. Several band ratios were applied to the target minerals and the results were bound to be satisfactory when compared to those of a HyMap reference survey. ASTER data was also used by Gad and Kusky (2007) for mapping lithology in Egypt. A 7-3-1 band combination was effective for distinguishing between granitic rocks and metamorphic rocks, while other band ratios successfully discriminated between vast rock units (gneiss, migmatite, amphibolite), volcanogenic sediments with banded iron formation, metabasalt and granitic rocks. The maps produced using ASTER data were

integrated with field observations to produce an enhanced geological map. In Namibia Gomez *et al.* (2005) successfully applied PCA to ASTER data in the visible near infrared (VNIR) and short-wave infrared (SWIR) spectra to improve geological mapping. Of the nine PCA components generated, the first five represented the most (98%) information, while the last four components were less useful as it contained abundant noise (Gomez *et al.*, 2005). The results provided new lithological boundary information that was used to upgrade the geological map of the area.

HyMap data was applied to map the Sarfartoq carbonatite complex in south-west Greenland (Bedini, 2009). The minimum noise fraction (MNF) transformation, a spectral reduction and data compression technique, was used. MNF is a form of PCA but instead of ordering the data by variance, the data is ordered according to a signal-to-noise ratio (Bertels et al., 2005; Chen, 2000). The first 20 MNF bands were used as input to an unsupervised classification algorithm to map carbonatite, carbonatite outer core zones (fenite with carbonatite dyke), fenite, marginal alteration zone (hematized gneiss) and gneiss. The classification achieved an overall accuracy of 87.6% and improved the geological mapping of the study area's carbonatite complex. HyMap and AVIRIS data were also used to map expansive clay soils in Colorado USA (Chabrillat et al., 2002). Smectite, illite and kaolinite were the targeted clay minerals. These clay minerals were identified by using higher-order MNF bands focusing on 2.0 to 2.45µm as clay absorbs wavelengths in this region of the spectrum. Field investigations gathered information on mineralogy and the swelling potential of the samples collected was determined by standard laboratory analyses. Together with EO images, maps of clay types with potential swelling were successfully compiled.

For many years EO for geological mapping was focused on the use of spectral information to detect alteration zones (Pour and Hashim, 2012a; 2012b) and for lithological mapping (Hewson *et al.*, 2005) without exploiting spatial information. It has been acknowledged that the incorporation of spatial information improves spectral classification (Blaschke, 2010), more so in geology due to its ability to infer geological transitions. Traditional EO lithological boundary detection is often subjective, with distinct variation occurring between expert analyses. Saliti *et al.* (2011) investigated the automatic detection of lithological boundaries based on the pattern rotational variant template edge matching algorithm of ASTER imagery in the South West of Iran. The accuracy of the detected lithological boundaries was spatially assessed by five experts against two reference sources: a published geological map produced by Iranian oil operation

companies (scale 1:100 000) and boundaries interpreted from a pair of stereo aerial photographs (scale 1:50 000). Although the results could not be quantified (due to no ground truth data of mineral composition), a comparison of the boundary detection with those of reference sources showed a spatial correspondence between rotational variant template edge matching technique and reference boundaries. It was noted that in heterogeneous areas this method did not correlate with the reference boundaries. Saha et al. (2011) applied an object-based approach to map drumlin landforms from 30m digital elevation model (DEM) in the Chautauqua drumlin field in North West Pennsylvania and New York. The shape and orientation of drumlin provides a wealth of information about the flow of past glaciers. Through the RGB combination of DEM, slope and aspect, drumlins were manually digitised by an expert and compared to the drumlins derived through automated GEOBIA multiresolution image segmentation algorithm. Using the drumlins which were manually digitised by an expert as a reference, the drumlins from automated method were visually and statistically compared to the reference drumlins. The results revealed that the automated procedure had over-identified the drumlins. Some of the over-identified drumlins were actually not drumlins, while other drumlins were underclassified by the manual method. The shape and length analysis of the drumlins that were comparable (overlaying each other) after the two methods (manual and automated) were applied, produced varied results, with the automated procedure having greater length values than the manual method. The need for focused research in GEOBIA, specifically for geological applications, can enhance the objectivity of results conducted by different expects.

5. EO image data fusion

Advances in sensor technology have significantly improved the spatial and spectral resolution coverage (from multispectral to hyperspectral) of remotely sensed data. Due to their respective capabilities and limitations, the different data types provide different representations of the same physical environment. When analysing an area it is desirable to use all the available data to increase the extraction of information and to improve the accuracy of image classifications (Zhang, 2010). To take full advantage of the increasing availability of EO data, image data fusion has been proposed (Houzelle and Giraudon, 1994; Khaleghi *et al.*, 2013; Maitre and Bloch, 1997; Pohl and Van Genderen, 1998; Simone *et al.*, 2002; Zhang, 2010). Image data fusion takes advantage of the capabilities

of each sensor and reduces its limitations. Acknowledging this, Khaleghi et al. (2013) maintain that the use of different sensors (or EO images) of the same scene provides a better understanding of the physical properties of the observed objects. Remotely sensed image data fusion is the integration of information acquired at different spatial and spectral resolutions from the same or different sensor platforms (e.g. satellite or aircraft) to produce fused data that contain more detailed information than a single image (Pohl and Van Genderen, 1998). Its advantage is the robust operational performance it provides, i.e. increased confidence, reduced ambiguity, improved detection and reliability, as well as enhanced classification ability (Du et al., 2013; Khaleghi et al., 2013; Pohl and Van Genderen, 1998; Zhang, 2010). EO data fusion can be performed on pixel/data, feature or decision levels. Pixel/data level fusion combines the original data from multiple sources into single-resolution data, whereas in feature level fusion the features of interest are extracted separately from the data and then combined into singlefeature datasets that may be processed further. GEOBIA has the unique ability to fuse disparate data sources based on groupings of pixels (objects), a capability that is not well researched, particularly in the geological domain.

The geological investigations such as the detection of potential faults, landslides and rock-fall have been conducted during the world's largest water conservancy project in China (Yang et al., 2000). Through the data fusion of SPOT panchromatic (2.5m) band together with SPOT multiresolution (10m) bands, a high spatial quality transformation of multispectral content was obtained. The second image fusion in this study was conducted with the enhanced SPOT multispectral fused with synthetic aperture radar (SAR) imagery. With the ability to detect sub-surface structures, together with high spatial multispectral resolution, the results from the image fusion improved the interpretation of linear structures in the study area. Pournamdari and Hashim (2014) also applied data image fusion of ASTER and Landsat TM for geological mapping of an ophiolite complex (the source of chromite ore deposits in the area) in South East Iran. A feature level fusion technique was applied on spectral angle mapper results derived from ASTER and Landsat imagery respectively to successfully derive a lithological map of the ophiolite complex.

Although data fusion enhances geological investigations, challenges still remain. The data imperfection, correlation, inconsistency and disparateness due to the diversity of the sensor technologies are among the data fusion glitches (Khaleghi *et al.*, 2013). While investigations have been conducted into data fusion algorithms (Du *et al.*, 2013; Pohl and

Van Genderen, 1998), no single data fusion algorithm is as yet able to overcome all the obstacles of data fusion. Research on data fusion suitable for geological investigations would eliminate inaccuracies resulting from the incorrect application of data fusion algorithms.

6. Conclusion

Technological developments in EO significantly influenced geological investigations. Recent improvements in sensor technology and the availability of higher spatial and spectral resolution data hold great potential for geological studies. EO techniques can be used to derive new information about the reflectance of rocks which can be employed for very accurate delineation of lithological formations. In some studies like urban monitoring applications, the challenges posed by increases in the intraclass spectra variability, triggered by increasing spatial resolution of imagery, can be addressed by the application of GEOBIA. More research on the use of GEOBIA for geological applications is urgently required. The recent launch of Sentinel 1a in April 2014, together with planned launch (in 2016) of Sentinel 1b RADAR imaging mission, provide potential for data fusion application with already launched Landsat 8 (February 2013), Worldview 3 (August 2014) and recently launched Sentinel 2a (and planned 2b) passive satellites. However, a number of challenges with data fusion have been noted. Most of these challenges arise from the data to be fused, imperfections and diversity of the sensor technologies, and the nature of the application environment (Houzelle and Giraudon, 1994; Khaleghi et al., 2013). A better understanding of these issues and the implementation of controlling measures will increase the potential of EO for geological applications. New technologies, such as GEOBIA, also offer potential solutions but more work is required to understand its value and implementation.

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