Incorporating a water-logging routine into CERES-Maize, and some preliminary evaluations

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Abstract

The inability of CERES-Maize v3.0 to simulate a fluctuating water table has been identified as a major constraint in using this particular model in South Africa and in Kenya. Information regarding fluctuating water tables under specific conditions, and their influence on maize production, has been presented in South African literature. The objective of this study was to construct a simple water-logging subroutine based on these mechanisms, using the existing input data structure of CERES-Maize. Results indicate that it is possible to simulate the impact of water logging on maize plants in this way. Further studies are needed to evaluate the assumptions made in this study and, if necessary, to make some refinements to the water-logging subroutine.

Introduction

In recent years it has become increasingly difficult to increase agricultural production by means of current methods of information transfer (Uehara, 1998). Most field trials conducted over the past decade have been aimed at refinement of recommendations made to the producer. It may be possible to stabilise African maize production if recommendations can be tailored to particular soil, plant and climatic conditions. In this respect the use of simulation models may serve to improve recommendations pertaining to variety selection, fertiliser use, irrigation scheduling, and optimum planting times (Acock, 1982; Thornton, 1990; Piper and Weiss, 1990).

The CERES-Maize model may be used for management decision making by the producer, provided that the model is evaluated and calibrated for various regions (Ritchie et al., 1998). This is an internationally-recognised maize model, highly acclaimed by researchers in the field (De Vos and Mallett, 1987; Du Pisani, 1987; Carberry et al., 1989; Piper and Weiss, 1990). The model was designed to use a minimum set of soil, weather, genetic and management information. It is a daily-incrementing model and therefore requires daily weather data consisting of maximum and minimum temperature, solar radiation and rainfall. It calculates crop phasic and morphological development using temperature, day length and genetic characteristics. Leaf area index, plant population and row width provide information for determining the amount of light interception, which is assumed to be proportional to biomass production. The biomass is partitioned into the various plant parts by means of a priority system. Water and nitrogen balance submodels provide feedback that influences the development of growth processes.

During a project to develop an integrated approach to assessing soil fertility and climatic interactions in pilot maize-producing areas in Kenya, it became evident that CERES-Maize v3.0 could not simulate the impact of water logging on maize growth (a relatively common occurrence in certain parts of Kenya), a conclusion supported by Hensley et al. (1997) and Hensley (2000).

Hensley et al. (1997) and Hensley (2000) describe the watertable logic for sandy soils as follows. There is an upward linear decrease in water content from the water table, where the soil water is at saturation, to the drained upper limit (DUL) of the layer approximately 600 mm above the water table. If the water-table depth is less than about 600 mm from the soil surface, aeration stress occurs for maize with a consequent decrease in yield. It can be expected that the number of live roots will decrease with an increase in time of water logging, and thus the uptake of nutrients will also decrease.

The objective of this study was to incorporate the Hensley logic into the water uptake and drainage subroutines of CERES-Maize. It simulates the rise and fall of the water table, with the concomitant effects of water logging on the growth of the maize plant.

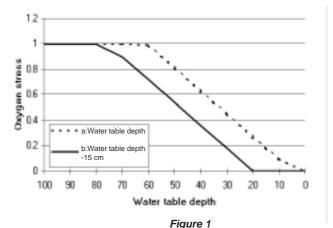
Material and methods

Crop growth simulation model

The CERES models for maize, sorghum, wheat, millet and barley were combined to provide a generic multi-crop model which runs with a single set of code, incorporating the development and growth sections for each individual model into a single module with a single soil component (Tsuji et al., 1994; Ritchie et al., 1998). According to Ritchie (1991), generic models should allow users to have more uniform procedures for validating models and for linking with components not included in the generic model. The generic CERES-Maize model (Hoogenboom et al., 1994), with modifications made by Du Toit (1996) to simulate secondary ears and tillering under low plant populations, was used for the simulations. An experimental approach to water-table simulation was used to simulate the results as reported in Hensley et al (1997). The results reported as "without water-table simulation" in this paper do not include the experimental approach.

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The relationship between water-table depth and oxygen stress

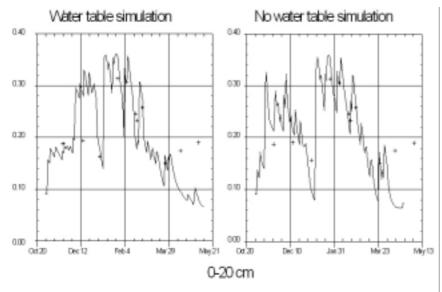


Figure 2

Measured (+) and simulated (-) volumetric soil-water content over time for the "with" and "without" water-table simulation for the 0 to 20 cm soil layer

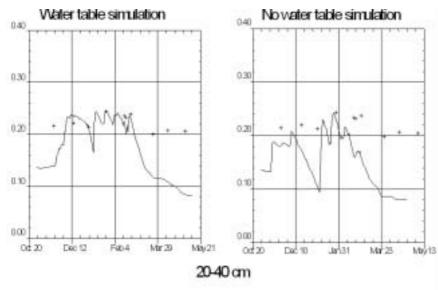


Figure 3

Measured (+) and simulated (-) volumetric soil-water content over time for the "with" and "without" water-table simulation for the 20 to 40 cm soil layer

Field trial

The trial was planted in the Mpumalanga Province of South Africa, in the Ermelo district. The climate is subhumid and the soil is hydromorphic (Longlands form, Ermelo family), with an Ehorizon at 600 mm. This horizon becomes water-logged during wet years, which occur frequently in the region. Water logging is the main constraint to maize production at this specific site. Rainfall during the maize growing season was above average at 735 mm. The cultivar PAN6479 was planted at 3.5 plants·m⁻² in 0.9 m row widths on 20 October 1993. Volumetric soil-water content was measured with a neutron hydro-probe at regular intervals during the growing season. The trial is described in more detail by Hensley et al. (1997).

Model modifications

The subroutine of CERES-Maize that simulates the draining of soil water (DRAINE.FOR) was modified to simulate the fluctuation of the water table and to calculate the oxygen stress index (OSTRESS). The following modifications where made to DRAINE.FOR. For CERES to identify if a specific soil has a tendency to form a water table or not, the saturation of the first soil layer, SAT(1), needs to be set to 0.40 as a volume fraction and the total drainage rate to 0.001 fraction ·d⁻¹. If either of these values are not met. no water-table simulation will occur. If the soil-water content of the first layer, SW(1), exceeds the drained upper limit of the same layer, DUL(1), all available water above DUL(1) will then drain (VLOEI) to the second layer. This is then repeated for consecutive layers, according to the "tipping bucket" concept of soil water modelling (Ritchie, 1998). In order for CERES to identify the depth of the water table (WTDIEP), the soilwater content of a specific layer, SW(L), needs to be greater or equal to SAT(L)-0.02. VLOEI is then changed from volume to a volumetric measurement by defining it with the soil depth and adding it to the soil-water content of the layer directly above the water table. If this layer SW(L) exceeds SAT(L)-0.02, it will become the WTDIEP and the process is repeated. If WTDIEP is less than or equal to 600 mm, oxygen stress OSTRESS is calculated as follows:

OSTRESS = 1.0-(AMIN1(1.0,(1.09 -0.018*(WTDIEP-15)))) (1)

AMIN1 is a Fortran 90 command - selecting the smallest value. The development Function 1 will be discussed later in this paper. In the growth subroutine GROSUB, OSTRESS was added to water and nitrogen stress to decrease growth. The source code for the modified DRAINE.FOR subroutine is shown in **Appendix 1**.

Results and discussion

Hensley et al. (1997) stated that if the water-table level rises above a soil depth of 600 mm, water logging would occur. To quantify OSTRESS, the assumption was made that at 600 mm depth no stress occurs, while increased stress will occur with each rising increment of the water table above 600 mm until complete saturation, and hence maximum stress occurs when the water-table level reaches the soil surface. Linear regression was therefore drawn between these two points.

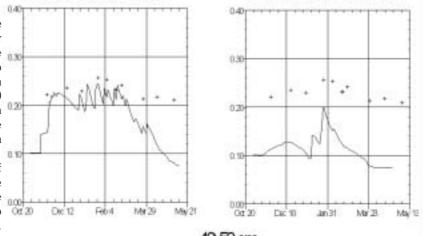
In GROSUB no stress is indicated by a value of 1 and full stress by 0; to use the same principle, the linear function was subtracted from unity. Because of an oxygen deficiency all roots can be expected to die in the soil layer just above the water table. Figure 1 indicates the influence of the water-table depth on oxygen stress without taking into account root die-back because of the influence on oxygen stress in the layer above the water table. To include this, 150 mm is subtracted from the water-table depth, as indicated in Function b of Fig. 1. This reaction to oxygen stress is simulated in the partially saturated soil above the water table. For this study a water-table depth is defined as the soil layer closest to the soil surface that is saturated. Further study is needed to find the correct function and to verify the different assumptions made in this study.

In the 0 to 200 mm layer (Fig. 2) the simulation line of the water table follows the same trends as the measured values (+). The point at 12 December indicates a marginal increase in soil-water content while the simulation shows a significant increase. This could possibly be explained by saying that the subsoil was not saturated at this stage. When the subsoil becomes saturated, the simulation accuracy increases.

At the end of the growing season the measured values increased while the simulated values decreased; this tendency appears in all the different layers (Figs. 3 to 7). One possible explanation is that CERES-Maize is over-predicting maturity and thus is still simulating water uptake even though measured soil-water content indicates that water uptake by the plant has stopped. The influence of water logging on the grain-filling period is in need of further study.

The degree of water saturation within a waterlogged soil is determined by the porosity which, in turn is related to the bulk density. In general it is calculated as 1-(bulk density/particle density) where mean particle density is assumed to be 2.65 g·cm⁻³. A volumetric water content of 0.40 will therefore be related to a bulk density of 1.60 g·cm⁻³. The particular E-horizon (500 - 700 mm) had a bulk density of 1.90 g·cm⁻³, a volumetric drained upper limit value of 0.20 and a saturation value of 0.28 (Hensley et al., 1997). The water content of the Ehorizon was most of the time for the season in the region of 0.24 volumetric content, indicating the highly saturated state and the subsequent waterlogging effect on the maize plant.

No water table simulation Water table simulation



50 cm

Figure 4

Measured (+) and simulated (-) volumetric soil-water content over time for the with" and "without" water table simulation for the 40 to 50 cm soil layer

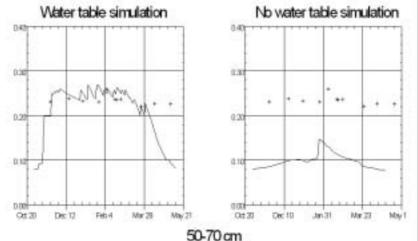


Figure 5

Measured (+) and simulated (-) volumetric soil-water content over time for the "with" and "without" water table simulation for the 50 to 70 cm soil layer

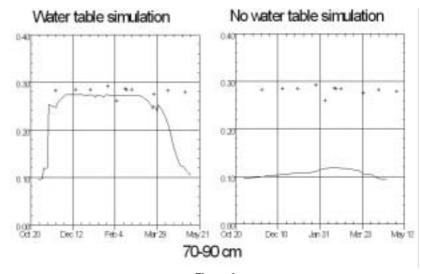


Figure 6 Measured (+) and simulated (-) volumetric soil-water content over time for the "with" and "without" water table simulation for the 70 to 90 cm soil layer

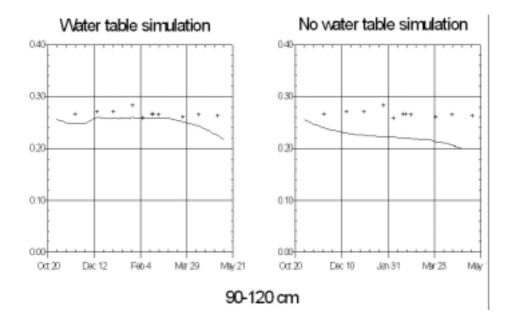


Figure 7 Measured (+) and simulated (-) volumetric soil-water content over time for the "with" and "without" water table simulation for the 90 to 120 cm soil layer

The measured yield for this trial ranging from 0 to 2 554 kg·ha⁻¹ with a simulated yield of 7 600 kg·ha⁻¹ that does not consider the effect of saturation, therefore simulating normal drainage. When the influence of oxygen stress (OSTRESS) was included, the simulated yield decreased to 2 480 kg·ha⁻¹ to give a much more realistic yield simulation.

A version of CERES-Maize v3.1 has now been released that has the ability to simulate tile drainage (Garrison et al., 1999). For the model to simulate water logging, the soil-input file has been modified to include depth and saturation values for each of the soil profile layers. While this information is available for this profile, it is lacking for the Kenya soils.

Conclusions

When the water table is less than 600 mm from the soil surface, oxygen stress has an influence on biomass production and it can have just as devastating an influence on maize yield production as water and nitrogen stress. With the simulation of water logging it is possible to quantify the influence of oxygen stress on maize yield production.

The advantage of this work is that it is now possible to simulate the fluctuation of the water table without additional soil inputs. It must be kept in mind that the DRAIN.FOR subroutine was extended only to accommodate an average saturated volumetric content of 0.40 (at an average bulk density of 1.60 g·cm⁻³). Although no provision was made to accommodate a series of saturation values, the empirical approach used showed significant improvement of yield simulation in water-logged soils. If high-quality soil input data are available, the more mechanistic approach to water-table simulation as represented in CERES-Maize v3.1 would be the preferred approach. However, if such data are not available, this more empirical water simulation could be used, and both approaches should be complementary to each other, depending on the quantity of available soil information.

The influence of oxygen stress on the kernel growth period needs to be investigated to improve the simulation of maize growth and development under water-logging conditions. With the development phase of the project completed the modified DRAINE.FOR subroutine needs to be validated against independent water-logging maize trials. The lack of such trials with high quality data is a constraint and trials that fit these criteria need to be executed.

Acknowledgements

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Appendix 1 Partial source code for the subroutine DRAINE.FOR
C***Simple approach to water table simulation.*****.
IF (SAT(1).EQ.0.40.AND.SWCON.EQ.0.01) THEN
IF (SW(1).GT.DUL(1)) THEN
VLOEI = (SW(1) - DUL(1))* DLAYR(1)
SW(2) = SW(2) + (VLOEI/DLAYR(2))
SW(1) = DUL(1)
VLOEI = 0.0
ENDIF
IF (SW(2).GT.DUL(2).AND.SW(3).LT.(SAT(3)-0.02)) THEN
VLOEI = (SW(2) - DUL(2)) * DLAYR(2)
SW(2) = DUL(2)
ENDIF
DO L= 1, NLAYR
WDIEP = WDIEP + DLAYR(L)
IF (SW(L).GE.(SAT(L)-0.02).AND.LAAG.EQ.0) THEN
LAAG = L
WTDIEP = WDIEP
ENDIF
IF (L.EQ.NLAYR.AND.LAAG.EQ.0) WTDIEP =WDIEP
END DO
IF (VLOEI.GT.0) THEN
L = LAAG - 1
SW(L) = SW(L) + (VLOEI/DLAYR(L))
IF (SW(L).GT.(SAT(L)-0.02)) THEN
VLOEI = (SW(L)-(SAT(L)-0.02))*DLAYR(L)
SW(L-1) = SW(L-1) + (VLOEI/DLAYR(L-1))
SW(L) = SAT(L) - 0.02
WTDIEP = WTDIEP - DLAYR(L-1)
IF (SW(L-1).GT.(SAT(L)-0.02)) THEN
VLOEI = (SW(L-1)-(SAT(L)-0.02))*DLAYR(L-1)
SW(L-2) = SW(L-2) + (VLOEI/DLAYR(L-2))
SW(L-1) = SAT(L-1) - 0.02
WTDIEP = WTDIEP - DLAYR(L-2)
ENDIF
IF (SW(L-2).GT.(SAT(L-2)-0.02)) THEN
VLOEI = (SW(L-2)-(SAT(L-2)-0.02))*DLAYR(L-2)
SW(L-3) = SW(L-3) + (VLOEI/DLAYR(L-3))
SW(L-2) = SAT(L-2) - 0.02
WTDIEP = WTDIEP - DLAYR(L-3)
ENDIF
ENDIF
IF (SW(3).GE.(SAT(3)-0.02)) THEN
VLOEI = (SW(3)-(SAT(3)-0.02))*DLAYR(3)
SW(2) = SW(2) + (VLOEI/DLAYR(2))
$SVV(\lambda) = SVV(\lambda) + (VLOEI/DLAIK(\lambda))$

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SW(3) = SAT(3) - 0.02
     WTDIEP = 30
  ENDIF
  IF (SW(2).GE.(SAT(2)-0.02)) THEN
     VLOEI = (SW(2)-(SAT(2)-0.02))*DLAYR(3)
     SW(1) = SW(1) + (VLOEI/DLAYR(2))
     SW(2) = SW(2) - 0.02
     WTDIEP = 15
  ENDIF
   ENDIF
C*****Determine water table depth***********
   WDIEP = 0
   LAAG = 0
   DO L= 1, NLAYR
   WDIEP = WDIEP + DLAYR(L)
С
     PRINT*, SW(L), SAT(L)-0.02, WDIEP, LAAG,L
   IF (SW(L).GE.(SAT(L)-0.02).AND.LAAG.EQ.0) THEN
    LAAG = L
    WTDIEP = WDIEP
   ENDIF
   IF (L.EQ.NLAYR.AND.LAAG.EQ.0) WTDIEP =WDIEP
   END DO
IF (WTDIEP.LE.60.AND.WTDIEP.GT.0) THEN
  OSTRESS = 1.0-(AMIN1(1.0,(1.09-0.018*(WTDIEP-15))))
  OSTRESS = AMAX1(0.05,OSTRESS)
   ELSE
  OSTRESS = 1.0
   ENDIF
  else
  OSTRESS = 1.0
  ENDIF
  VLOEI = 0
  LAAG = 0
  WDIEP = 0
  WTDIEP = 0
  OSTRESS = AMAX1(OSTRESS,0.01)
```